Toward monitoring ocean wave activity using seismic stations



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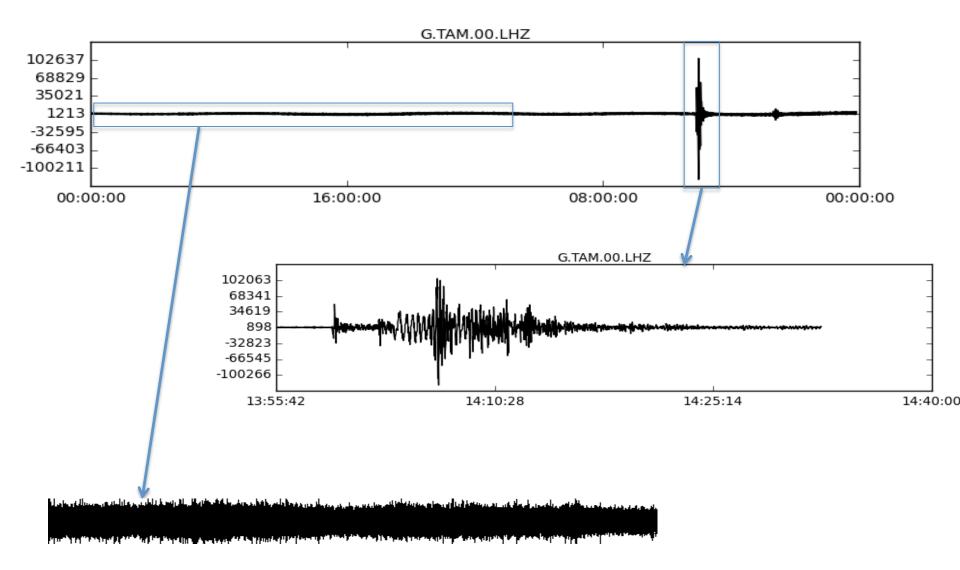








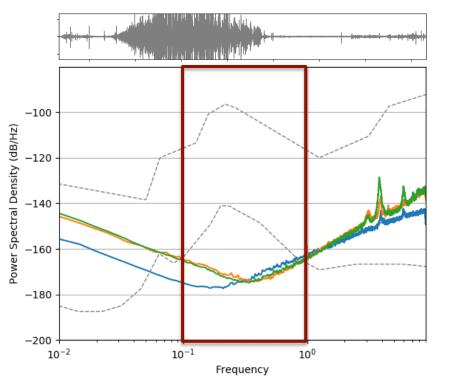
Greece - Magnitude:6.1



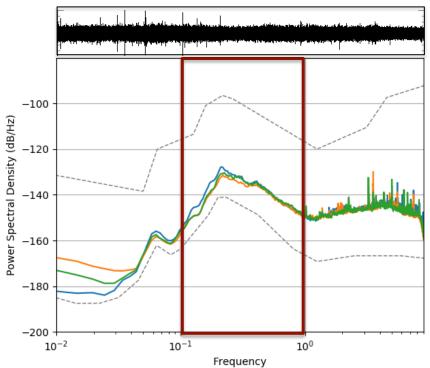




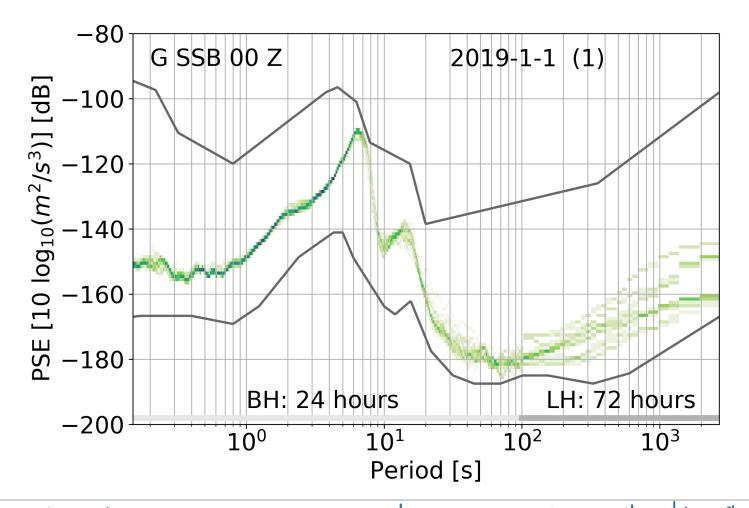
MARS



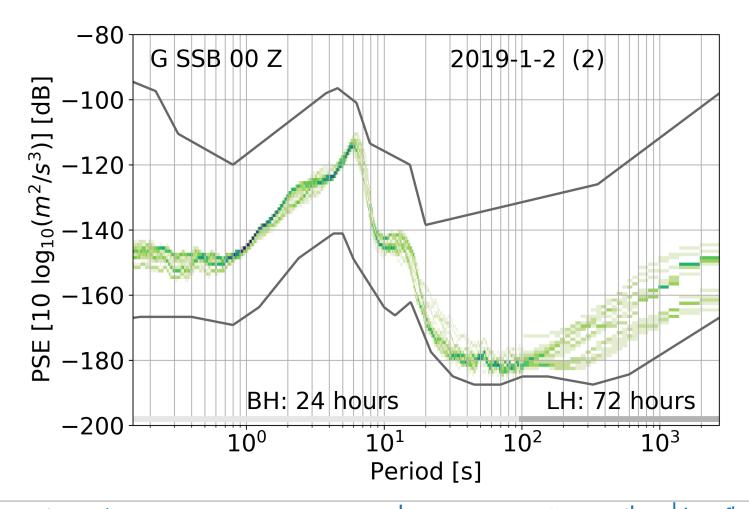
EARTH



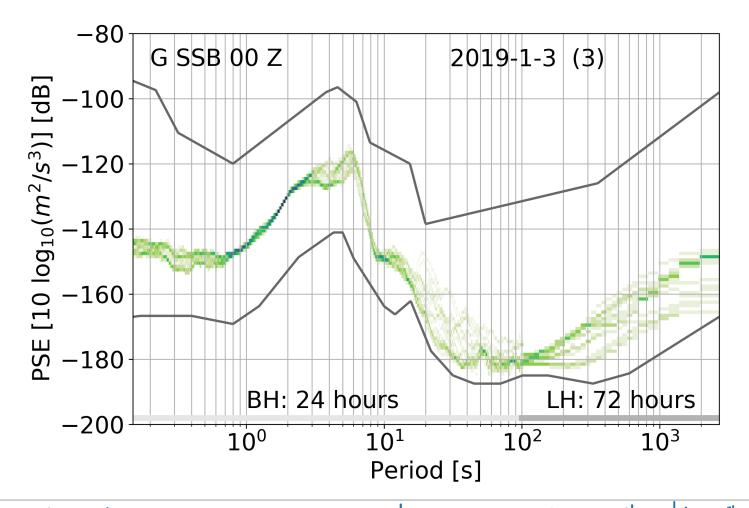
No ocean on Mars \rightarrow No microseisms \rightarrow Noise level 1000 lower between 0.1-1 Hz



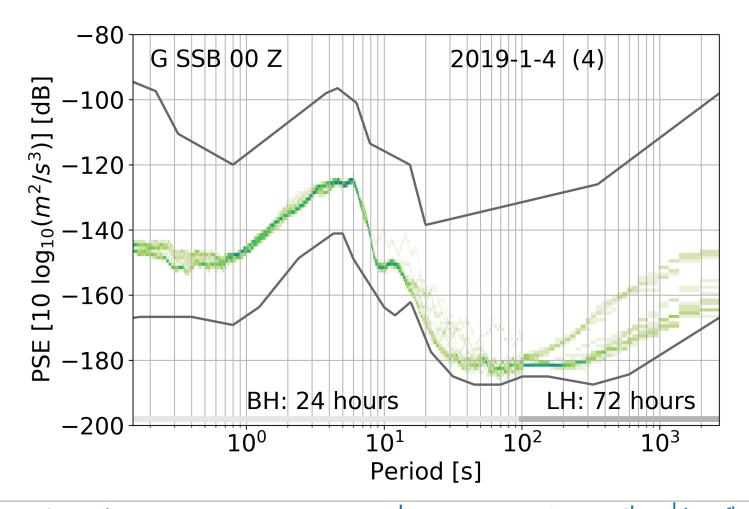
والمراوية والمساورة والمراك والمراحية والمرجوع والمراجع والم والمراجع والمراجع والمراجع والمراجع والمراجع والم والمراجع والمراج



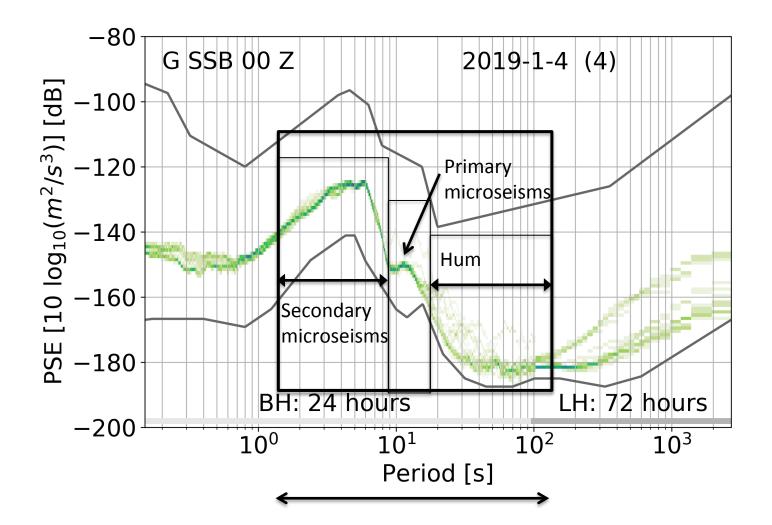
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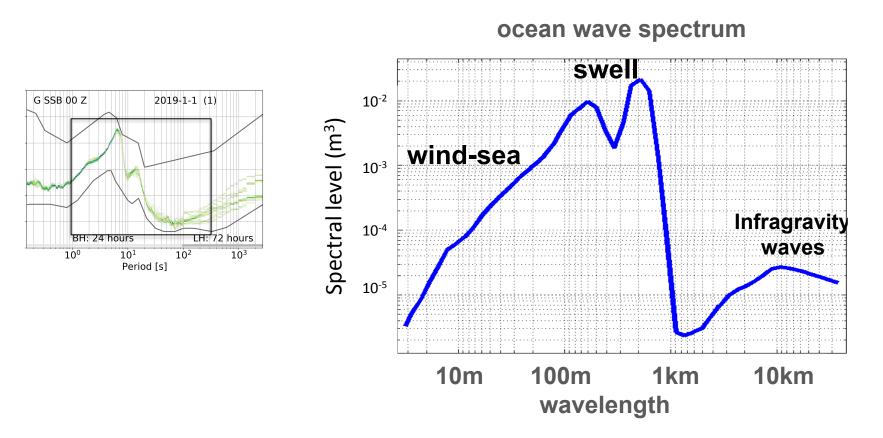


والمراوية والمحاد والمراك والمراحية والمرجوع والمراحل والمرجوع والمراجع والمراجع والمرجوع والمراجع والمرجوع والمرجوع والمراجع والمرجوع والمراجع والمرجوع والمراجع والمرجوع والمراجع والمرجوع والمرجوع



Noise with periods between 1 and 500 sec are generated by oceans waves

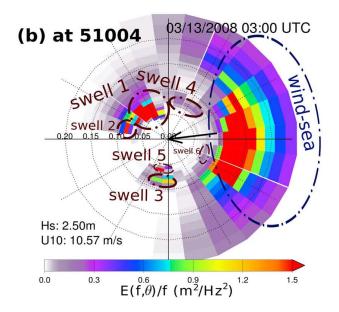
Ocean wave spectrum



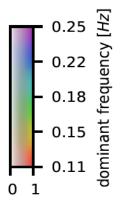
Wind sea and swell \rightarrow primary and secondary microseisms Infragravity waves \rightarrow hum

> Longuet Higgins, 1950 Hassselman 1963 Ardhuin, Gualtieri, Stutzmann, 2015

Secondary microseisms (period 1-10 s)



a) Pressure Model

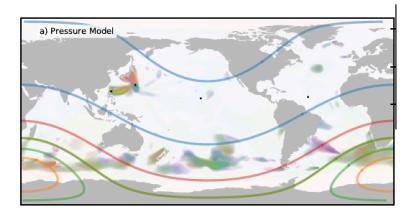


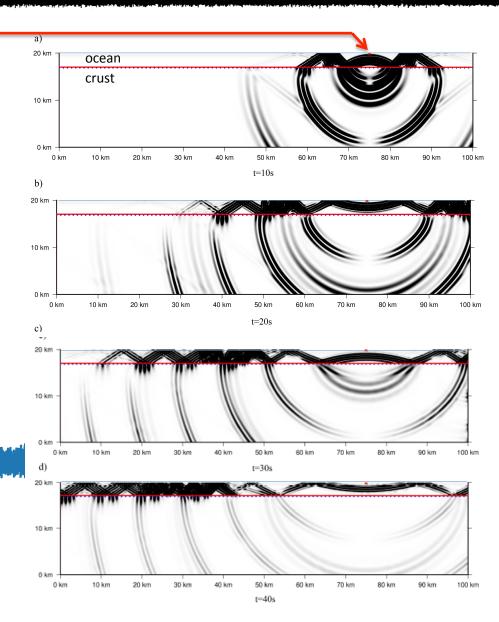
Waves are computed every 6 hours Code WAVEWATCH III 6-hourly wind analysis from ECMWF

Ardhuin et al., 2011

Secondary microseisms (period 1-10 s)

Pressure sources every 3 hours (IFREMER model)





large amplitude surface waves and tiny body waves

Surface waves: Rayleigh waves source site effect

30°N

30°S

60°S

180

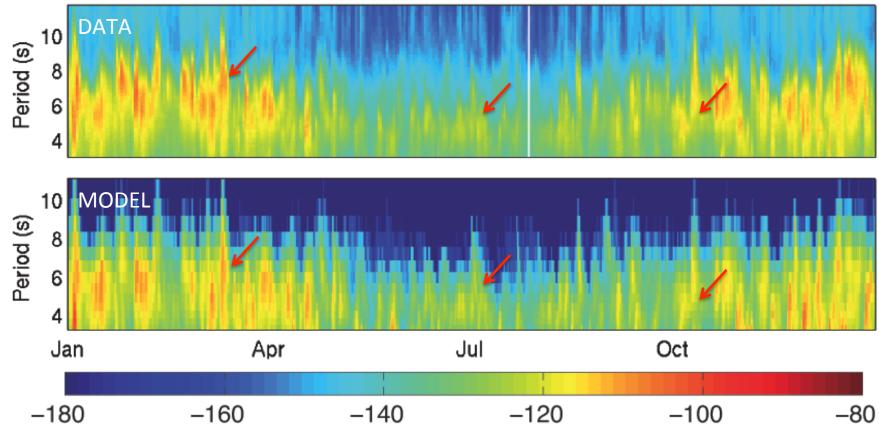
Amplification factor for the seismic wave period T=6s 0.8 0.7 60⁰N 0.6 30⁰N 0.5 0.4 00 Period=6s 0.3 30⁰S 0.2 60⁰S Bathymetry 0.1 1000 n 120⁰W $60^{\circ}W$ 60⁰E 120⁰E 180⁰W 180⁰W -2000 -3000 Amplification factor for the seismic wave period T=10s -4000 0.8 -5000 0.7 6000 60⁰N 7000 DRV 0.6 20°E 180°W 30⁰N 0.5 0.4 Period=10s 0.3 30⁰S 0.2 60⁰S 0.1 0 120⁰W 60⁰W 60⁰E 120⁰E 180°W 180°W 00

Longuet Higgins, 1950, Kedar et al., 2007, Ardhuin et al., 2011, Stutzmann et al., 2012



Seconday microseisms surface waves

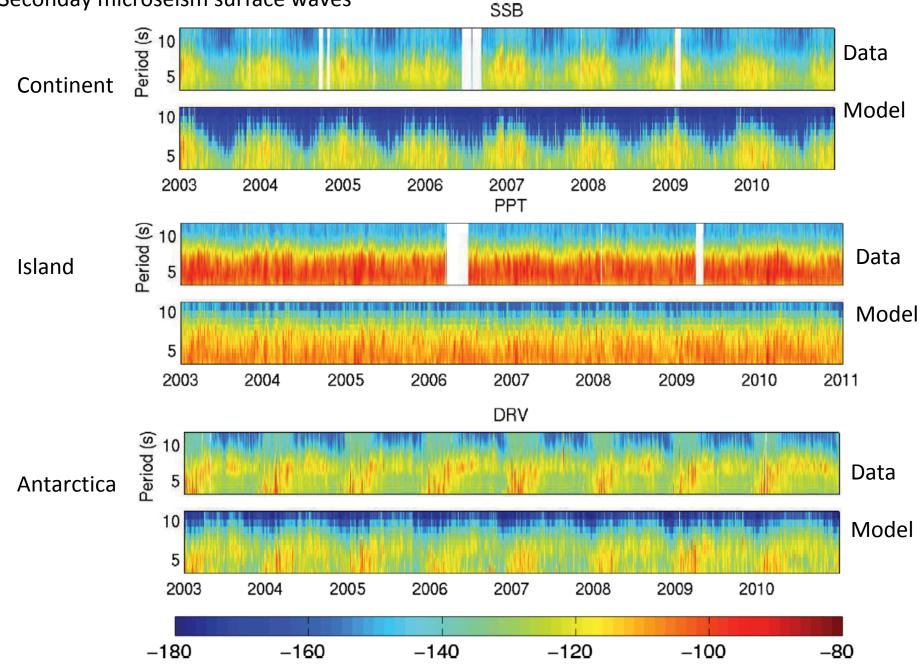
SSB 2008



Spectrogram are well modelled (frequency content and amplitudes) Strongest PSD are due to large storms

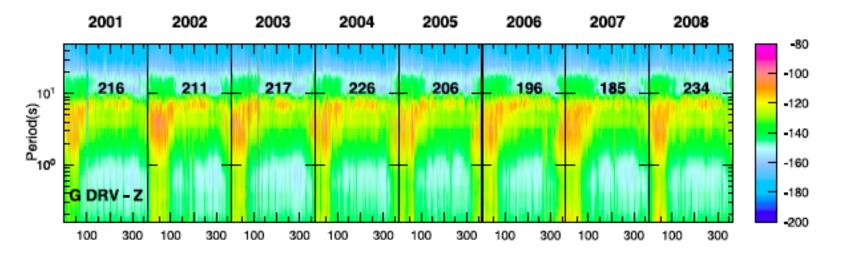
Weaker PSD is due to coastal sources related to ocean wave coastal reflection

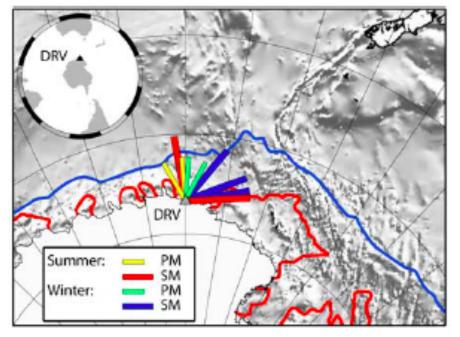
Seconday microseism surface waves



Stutzmann, Ardhuin, et al., GJI, 2012

Sea ice effect on seismic noise

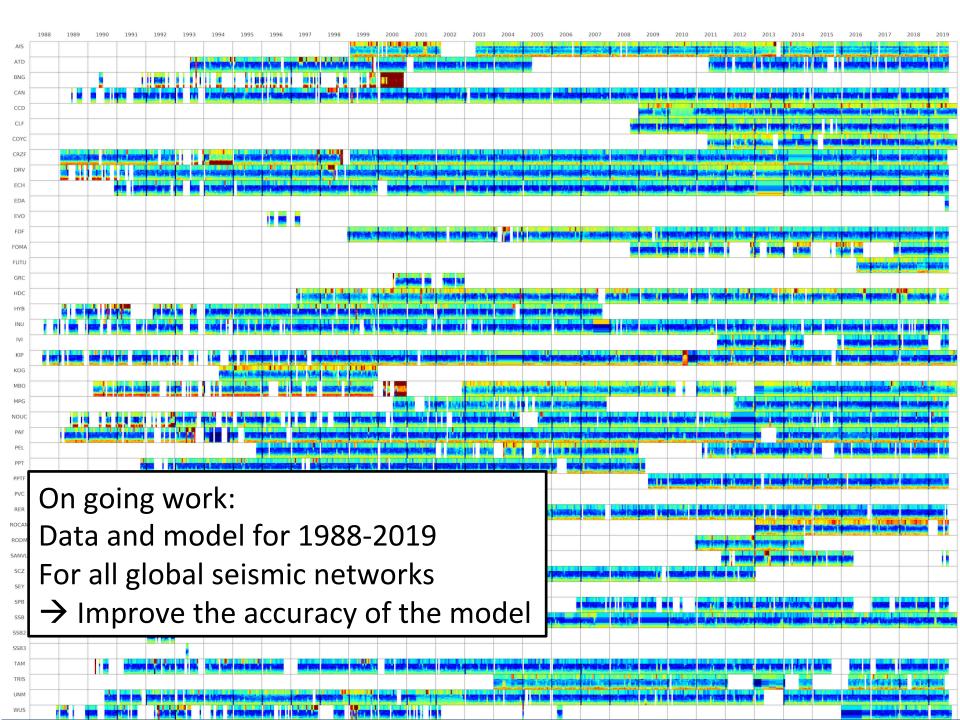




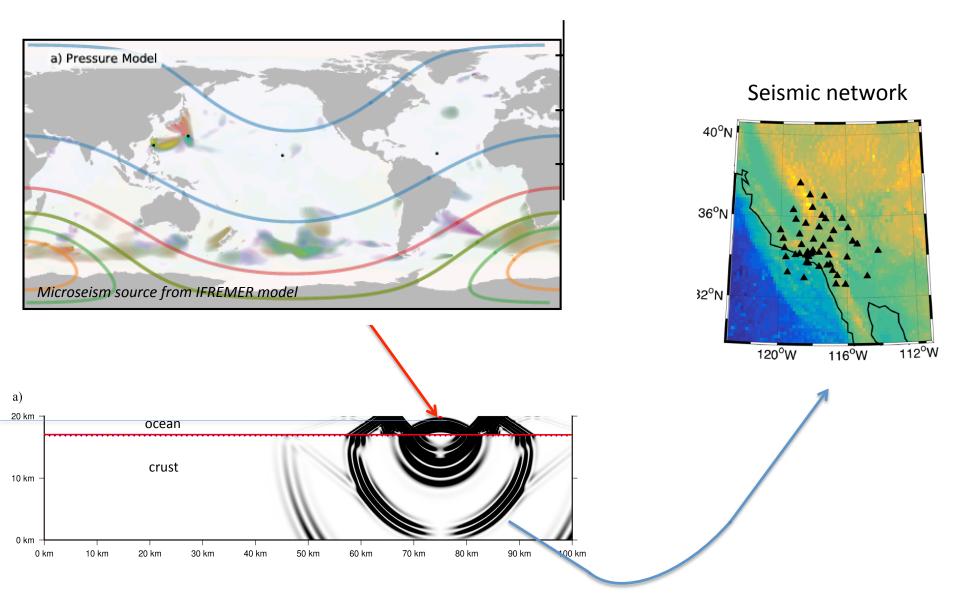
In winter, decrease of the amplitude of

- the primary microseism (10-15sec)
- the short period secondary microseism

Stutzmann, Schimmel et al., 2009 ; Grob et al., 2011

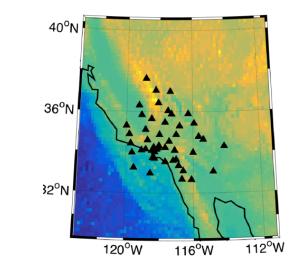


Secondary microseism body waves



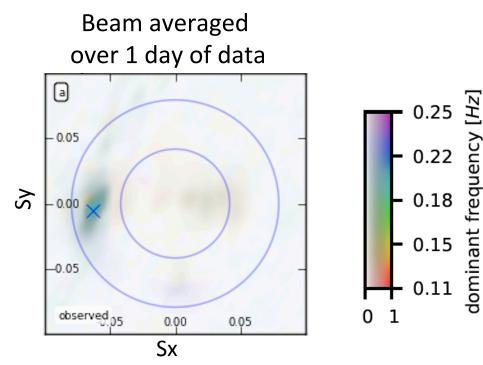
Secondary microseisms body waves detection

dominant frequency

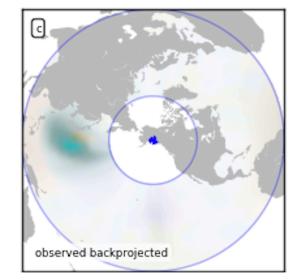


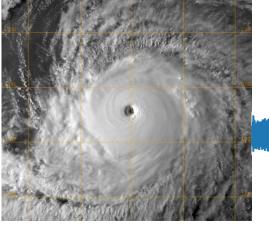
The array record P-waves from multiple sources

Each source is defined by its location, corresponding to a P-wave slowness: $s = (s_x, s_y)$ and its dominant frequency f

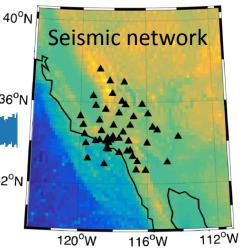


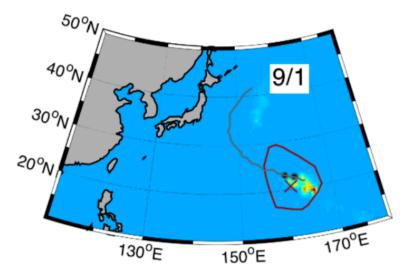
Back projection

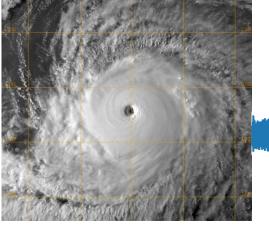




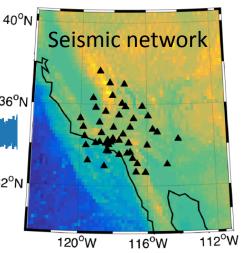


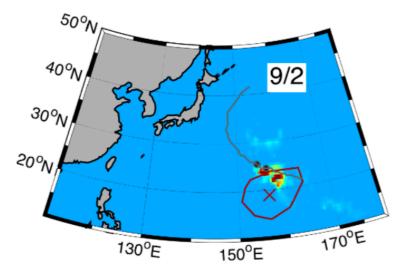


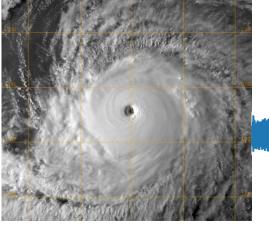




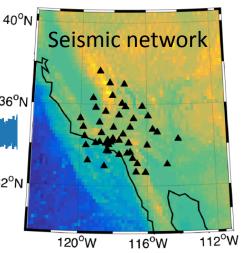


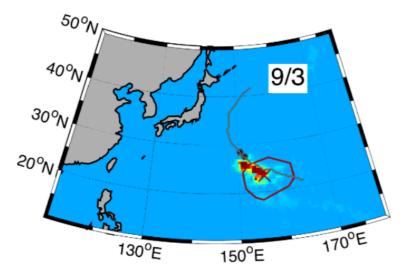


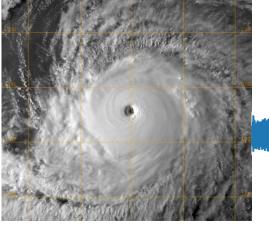




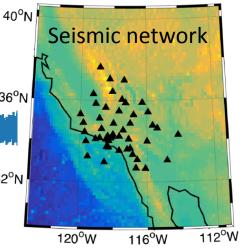


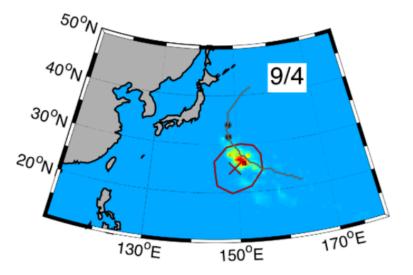


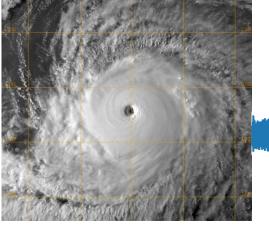




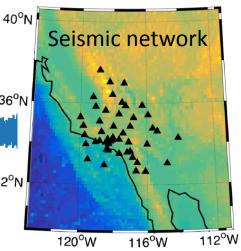


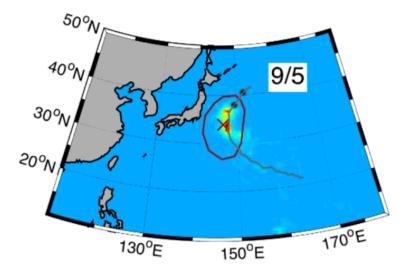




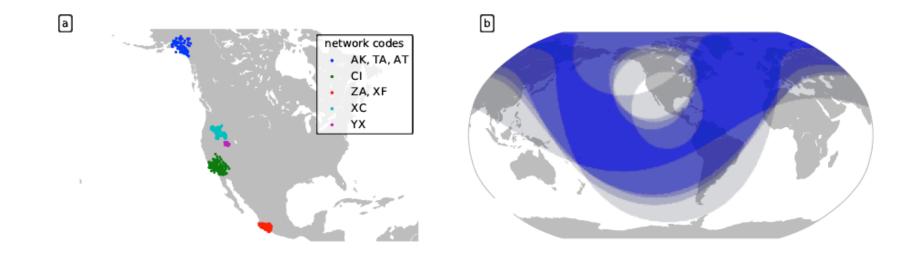


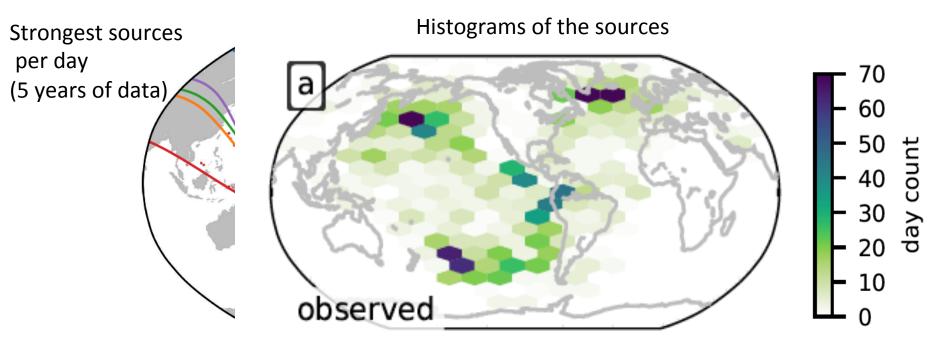




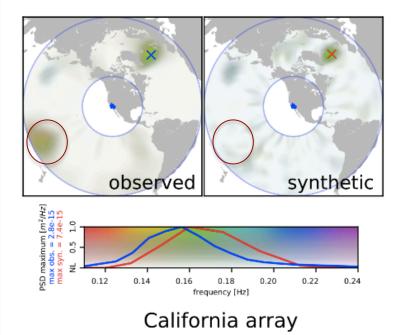


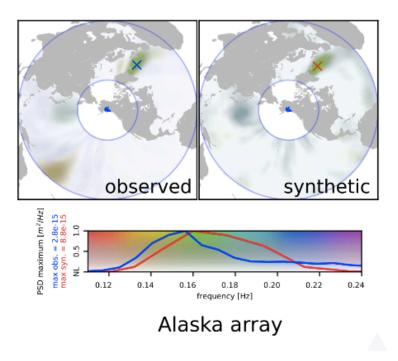
Secondary microseism sources at global scale

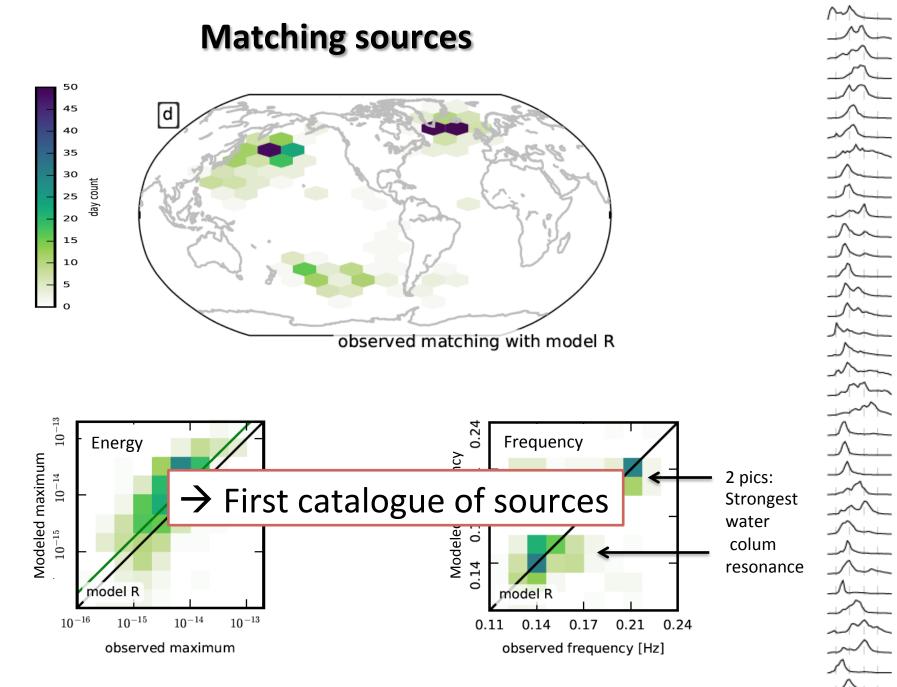




Comparison of back projected sources



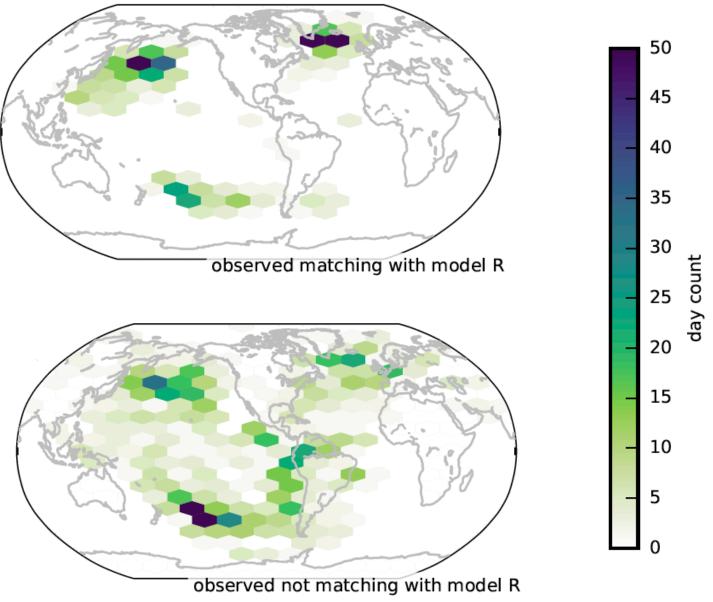




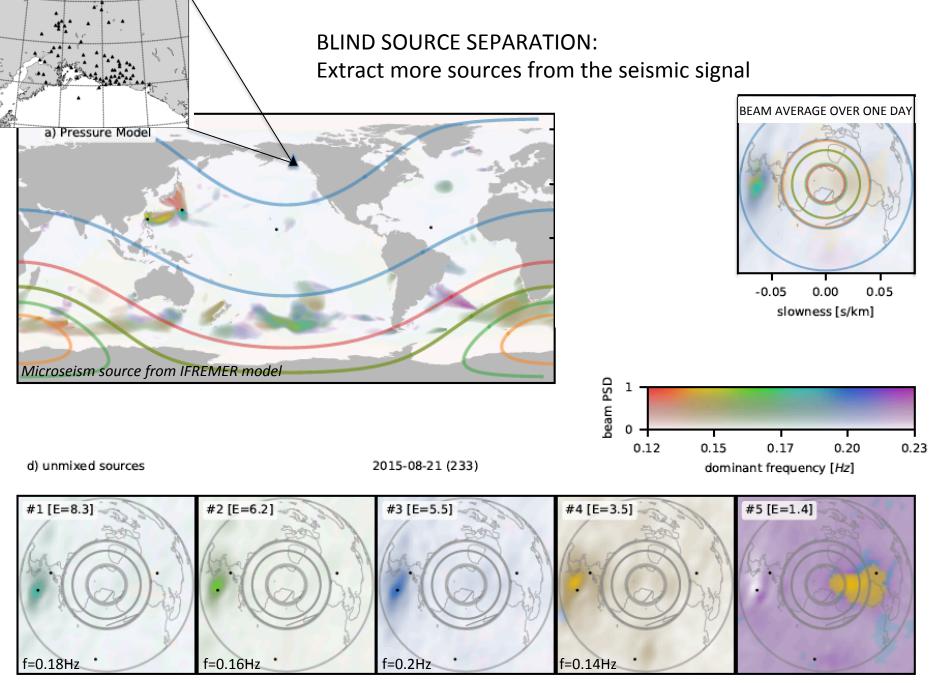
Energy and dominant frequency are accurately modeled

Meschede, Stutzmann et al., JGR, 2017

All sources



Meschede, Stutzmann et al., JGR, 2017



Meschede, Stutzmann, Schimmel, 2019

Conclusions

- Seismic data provide long time series that can be accurately modeled using sources from oceanographic models
- Body waves enables to extract individual sources
- Seismic data are very sensitive to the wave coastal reflection coefficient

On going work:

- Analysis of longer time series
- Machine learning for building new catalogue of sources
- Improve the modeling of noise

